

A study of dynamical processes in the middle atmosphere associated with a sudden stratospheric warming using a high-resolution and high-top general circulation model

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Dynamical events called stratospheric sudden warmings (SSWs) greatly alter the thermal and dynamical conditions in the whole middle atmosphere. In past several SSW events, it has been reported that the stratopause once disappears after warmings and reforms at an elevated height [e.g., Siskind et al., 2007; Manney et al., 2008, 2009], which is called the “elevated stratopause” (ES). Tomikawa et al. [2012] showed that the ES forms and descends to the climatological height of the stratosphere due to the westward gravity wave (GW) forcing (GWF), while some studies indicated that planetary wave (PW) forcing (PWF) plays a key role in the initial formation of the ES [e.g., Limpasuvan et al., 2012, 2016]. Despite these intensive researches, the relative contributions of GWs and PWs to the ES formation remain to be elucidated. Yamashita et al. [2013] suggested the importance of the lateral propagation of GWs [Sato et al., 2009, 2012] in the ES formation. To elucidate the mechanisms of the dynamical processes associated with SSWs such as the ES, we conduct a quantitative study using the outputs of the simulation of the 2018/2019 SSW event carried out by a GW-permitting general circulation model that covers from the troposphere to the mesosphere and lower thermosphere (MLT). A quantitative and three-dimensional analysis is conducted from the point of view of the interplay of GWs and PWs in the middle atmosphere [Sato and Nomoto, 2015].

In the 2018/2019 event, the ES forms around 10 January, 2019 after the major warming on 1 January (Fig. 1). In addition, it is observed that a significant temperature maximum appears in the polar upper mesosphere around 28 December prior to the disappearance of the stratopause, which is referred to as the “double stratopauses” (DS). Then we focused on the formation of the DS and ES and analyzed the role of wave forcing in these phenomena.

First, before the DS formation, maxima of static stability N^2 and potential vorticity (PV) are formed in the regions between positive and negative GWF. Then $PWF > 0$ is observed on the polar side of this PV peak. It is suggested that this PWF acts to eliminate the anomalous PV gradient. Subsequently, $PWF < 0$ appears in the polar MLT and the DS is formed at $z \sim 85$ km. Therefore, plausible processes of the DS formation are that 1) the residual mean vertical wind driven by GWF induces the formation of the PV maxima due to increase of N^2 , 2) the barotropic (BT)/baroclinic (BC) instability occurs on the polar side of the PV maxima, 3) and then PWs are generated and exert forcing in the polar upper mesosphere which induces downwelling and the associated temperature maximum (Fig. 2).

Second, on 10 January when the ES formed, both GWF and PWF are strongly negative at $z > 80$ km in the winter polar region. Therefore, it is indicated that these wave forcings are responsible for the ES formation. The eastward jet is reformed in the polar mesosphere on that day. This fact suggest that the eastward jet enables the GW and PW propagation into the polar upper mesosphere. The eastward jet recovery occurs mainly due to the radiative cooling in the polar night region. It is considered that the latitudinal gradient of the temperature and thus the height of the recovered jet and the ES are significantly affected by the temperature structure associated with the DS.

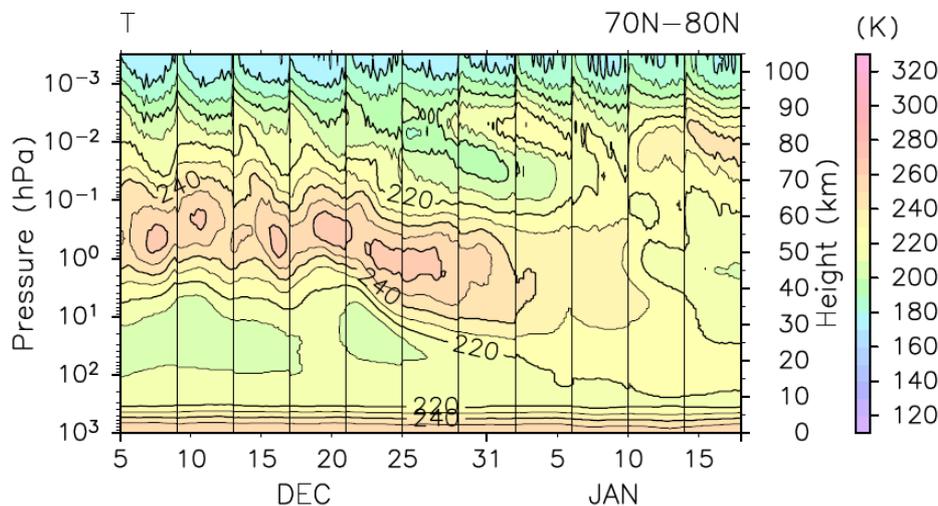


Figure 1 Time-height sections of zonal and 70°N – 80°N regional mean temperature. The vertical lines represent the boundaries of the model runs.

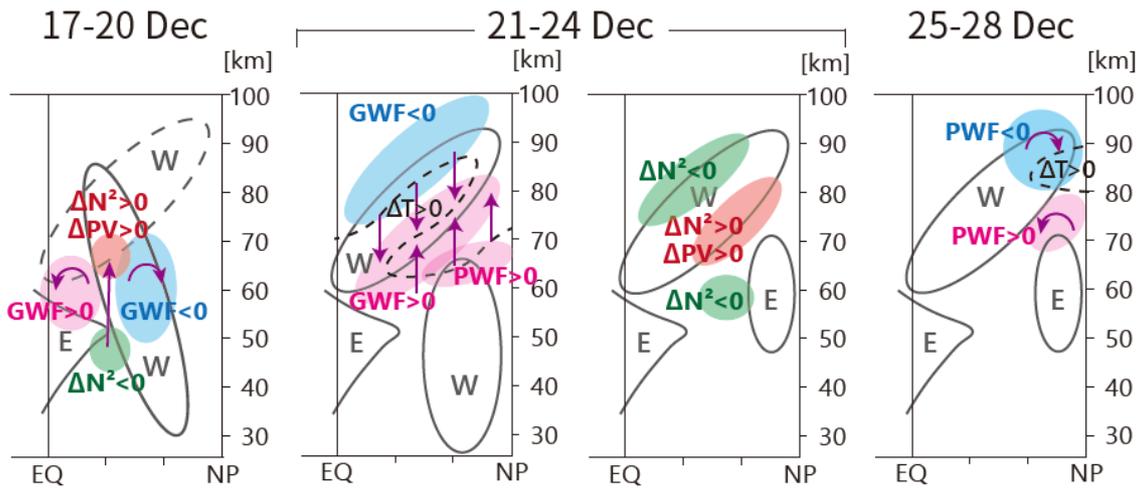


Figure 2 A schematic illustration of the formation mechanism of the DS in the meridional sections. “W” and “E” denotes westerly and easterly winds. Δ means the anomaly of each value. The purple vectors indicate the direction of residual mean circulation.

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